

Nanodiamonds do not provide unique evidence for a Younger Dryas impact

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Microstructural, $\delta^{13}\text{C}$ isotope and C/N ratio investigations were conducted on excavated material from the black Younger Dryas boundary in Lommel, Belgium, aiming for a characterisation of the carbon content and structures. Cubic diamond nanoparticles are found in large numbers. The larger ones with diameters around or above 10 nm often exhibit single or multiple twins. The smaller ones around 5 nm in diameter are mostly defect-free. Also larger flake-like particles, around 100 nm in lateral dimension, with a cubic diamond structure are observed as well as large carbon onion structures. The combination of these characteristics does not yield unique evidence for an exogenic impact related to the investigated layer.

transmission electron microscopy | diffraction | spectroscopy

Recently, the presence of different types of diamond morphologies and structures in a black carbon-rich sedimentary layer has been used as evidence to support an extraterrestrial impact approximately 12.9 ka ago as the cause of the Younger Dryas climate change (1–3). These authors argue that shock metamorphism produced these diamond particles (2). Other arguments from the same authors supporting the suggestion of an extraterrestrial impact at the Younger Dryas are, amongst others, high concentrations of Ir, magnetic microspherules and grains and carbon spherules. However, strong doubts still exist about the nature and implications of these markers. Indeed, Paquay et al. (4) in 2009 failed to reproduce the elevated concentrations of platinum group elements (including Ir) reported by Firestone et al. (3). Moreover, the same authors showed that $^{187}\text{Os}/^{185}\text{Os}$ isotopic ratios, considered to be the most sensitive indicator of a meteoritic contribution, yield values typical of the terrestrial crust. Surovell et al. (5) also failed to confirm enhanced levels of impact microspherules in the Younger Dryas unit from seven different sites. Also, the recent electron microscopy data presented to prove the existence of Lonsdaleite particles with crystallographic defects, claimed to be due to impact-related shock transformation, is not convincing (2). Although the reported lattice parameters can indeed be assumed to originate from Lonsdaleite, low precision measurements from ring patterns resulting from small particles can at most be seen as a first indication. This uncertainty was recently confirmed by an extensive study by Daulton et al. (6) who compared the Kennet et al. electron diffraction data (2) with different graphene morphologies and simulated Lonsdaleite diffraction. The lack of any experimental intensity at expected Lonsdaleite lines such as 102 indicates that no such structure appears in the investigated samples. In fact, the occurrence of asymmetric double diffraction lines points to the existence of graphene/graphane aggregates, also explaining all other observed rings. Also, no two-beam or high-resolution lattice images were presented by Kennet et al. to support their interpretation of stacking faults in the Lonsdaleite crystals.

In the present work, we focus on diamond particles and carbon isotope values in material obtained from excavations in Lommel (Belgium), in which the Younger Dryas unit or boundary (YDB) is revealed as a thin black layer as shown in Fig. 1. The same layer was also included in the original paper claiming the extraterres-



Fig. 1. Depth location of the black Younger Dryas Boundary at Lommel, Belgium.

trial impact (3) and in the recent confronting study of Paquay et al. (4). The thickness and depth of this black layer vary depending on the actual location investigated along the excavated trench. At most places, the layer is buried under approximately 50 cm of yellow clay whereas 1 m thick, fine white sands occur underneath. Samples of this black layer essentially composed of carbon, as well as samples from parts of the immediately over- and underlying sediments, were investigated by different modes of transmission electron microscopy (TEM) and $\delta^{13}\text{C}$ isotope measurements to document the crystallography and possibly the formation process of the diamond particles as well as the origin of the C matter present in this layer.

Results

Our findings confirm, and in fact reveal more direct proof than the earlier studies, the existence of diamond nanoparticles also in this European YDB layer. No such particles are found in the overlying silt and clay or in the underlying fine sands. The latter contain, aside from the prevalent amorphous carbon material, the expected silicates and calcites with some of those also occasionally observed in samples of the black layer. In contrast to the reports from other YDB sites (3) the Lommel black layer samples did not contain any millimeter-sized foam-like carbon spherules. Because such spherules are relatively brittle due to a high degree of porosity of the carbon skeleton, a possible reason for the lack of such large particles in the Lommel layer could be the action of diagenesis through which the particles are crushed under the weight of the upper soil layers. On the other hand, any existing diamond material in such spherules, as found before in carbon spherules collected in undisturbed upper soils (7) from Belgium and Germany, would still be present in the Lommel black layer.

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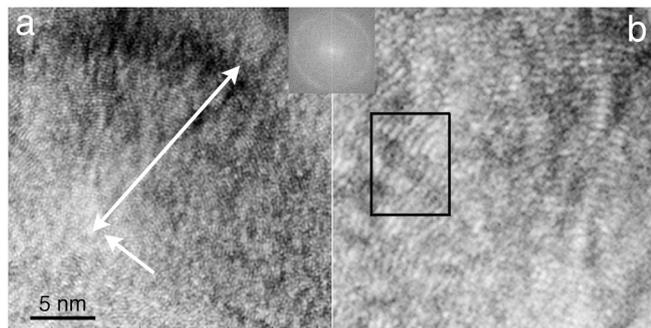


Fig. 5. (A) Carbon onion with compressed interplanar spacings between 0.268 nm and 0.284 nm [with magnified part and highlight of defect region for clarity in (B)]. The single arrow indicates the center of the onion and the double arrow its radius. The inset is an FFT revealing the close-to-circular symmetry and sharp interlayer distance.

such carbon onions can act as nanoscopic pressure cells for the formation of diamond nanoparticles in the core of the onion structure (13, 14). Due to the size of the onions yielding severe overlapping of above and below curved sheets plus the surrounding amorphous carbon, the images are unfortunately not sufficiently clear for any further structural conclusion, even after applying image processing.

In view of earlier reports of Lonsdaleite with respect to impact-related material, we deliberately looked for hexagonal diffraction patterns with the proper lattice parameters in our samples. However, no indications of such a structure were found, not in the present material from Lommel [nor in some revisited samples from the spherules collected from undisturbed upper soils (7)]. In these experiments special care was taken to avoid any effects of the high voltage electron beam in the TEM, which might damage or even transform the Lonsdaleite back into graphite due to the instability of this structure.

Also carbon isotope measurements and C/N values were determined from the black material of the Lommel YDB layer. The nanodiamond particles in the present material could not be analyzed separately because of their small size. The Lommel samples were thus analyzed in aliquots of increasingly smaller and smaller size fractions (down to $\sim 10 \mu\text{g C}$) in an attempt to eventually detect an effect driven by the presence of extraterrestrial organic material. The $\delta^{13}\text{C}$ and C/N values remain constant through the experiment and generated $\delta^{13}\text{C}$ values between -28.1 and -26.3‰ , and C/N ratios between 9.2 and 26.9.

Discussion

Comparing the findings of the present work with those in the top-soil spherules (7), those from Kennett et al. (1, 2) and other meteorite findings (15–18), leads to the following remarks. Both micron-sized flakes and nanoparticles of cubic diamond are present in the Lommel YDB material and the top-soil spherules. However, there is no report of flakes or excessively thin cubic diamond in meteoritic or impact diamonds. The smallest nanoparticles, up to a few nm in diameter, do not contain any one- or two-dimensional crystal defects in either the Lommel or top-soil cases. Larger sized nanoparticles, around 20 nm diameter, revealing multiple twinning were only observed in the Lommel samples, whereas in the top-soil spherules the larger agglomerates consist of nanostructured grains with strongly deformed crystal lattices (7). Moreover, most twins in the Lommel particles appear in a nonlinear fashion, even star-like, indicating a close-to isotropic type of growth. The larger flakes, around 100 nm in lateral dimension, do not reveal any twins. Stacking faults were not observed in either of our samples. Kennett et al. (1, 2) reported no twinned particles or flakes. The details of the observations on the diamond in the carbon spherules from undisturbed top-soils were attrib-

ted to a CVD type growth process, rather than to explosion related phenomena (7).

A very comprehensive work on twinned diamond nanoparticles from meteoritic material was performed by Daulton et al. (15) who compared artificial shock-synthesized (by detonation) and low-pressure chemical vapour deposition (CVD) material with residues from the primitive Allende and Murchison carbonaceous meteorites. One important conclusion was the similarity between the CVD and meteoritic material in view of the number and type of twinned particles. In the present study, most larger particles [i.e., with diameters above 10 nm (but excluding the flakes)], show some kind of twinning (mostly five-fold star, some parallel), the smaller ones being defect-free. Especially the dominant nonlinear character of the observed twins, points to an isotropic growth mechanism for the nanoparticles, possibly of the CVD type in contrast to the detonation type.

In the same work, Daulton et al. (15) also looked at dislocations, but did not find any in the meteoritic or CVD samples. The shock-synthesized material, however, did reveal clear signs of dislocation defects, be it in nano-particles instead of in flakes of the size observed in the Lommel samples. Such dislocations were interpreted as resulting from a martensitic-like transformation from graphite to diamond (19). This could imply a shock-related history for the flakes, where their microscopic shapes could be seen as a remnant of a shock wave front passing through the material. Although the latter is indeed a possible and likely interpretation in the case of particles of known origin, other potential pathways to produce dislocations in the flakes could easily be envisaged, including lattice misfits during epitaxial deposition and growth and strain hardening during deformation. Moreover, more recent work on nanodiamond from detonation experiments also did not reveal any flake-like shapes (20), whereas the type of twinning confirmed the observations by Daulton et al. (15).

Double-check of both the present material and the top-soil spherules for crystallites with a hexagonal symmetry did not reveal Lonsdaleite in either of those. According to Daulton et al. (15), Lonsdaleite nanocrystals are relatively stable in the electron beam, so little or no back-transformation to graphite during observation is expected. The YDB material reported on by Kennett et al. revealed cubic diamond nano- as well as microparticles. In the most recent report by Kennett et al. the existence of Lonsdaleite was identified on the basis of electron diffraction patterns (2). The images of these particles seem to correspond to the flake morphology. However, recent work by Daulton et al. (6) clearly shows that the diffraction rings observed by Kennett et al. (2) can also be explained by graphene/graphane aggregates whereas a 102 ring expected for Lonsdaleite was not observed. Kennett et al. also claimed the observation of stacking faults, but the experimental evidence cannot be considered as very convincing. From the lack of Lonsdaleite in our present material and the results from Daulton et al. (6, 15) it is concluded that no shock-induced metamorphism has occurred when the Lommel material was formed. Alternatively, the flake shape could be a remnant of a deposition and delamination process during which some dislocations are formed by the respective mechanisms proposed above.

The carbon onions observed in the Lommel material are relatively large and can consist of around 50 graphene sheets yielding a diameter of around 25 nm and containing several stacking defects in their shell structure. In most other cases where carbon onions have been observed, they usually have a much smaller diameter and consist of only a few graphene sheets depending on the processing history. An example hereof are the carbon onion structures with diameters below 10 nm observed in artificially produced wood charcoal (21). The latter were attributed to the carbonization of lignin structures in the raw wood starting material. Larger rounded graphitic structures have recently been observed in modern and fossilized charcoal structures indicating a large preservation time after a formation during natural wild

could indicate a martensitic-like transformation path from graphite to diamond, although epitaxial growth cannot be ruled out. Because no hexagonal diamond (Lonsdaleite) could be identified a shock-induced mechanism is unlikely to be involved in the formation of the crystalline carbon material in the present layer. The observed large carbon onion structures could be produced by a variety of processes, which implies that their exact origin remains unclear. Graphitic ring structures can indeed be formed during natural processes such as wild fires, although the existing observations reveal slightly different morphologies. As a final conclusion it should be stated that the present variety of crystalline structures observed in the black Younger Dryas boundary in Lommel does not provide sufficient evidence to conclude an exogenic impact as the origin of these structures.

- Kennett DJ, et al. (2009) Nanodiamonds in the Younger Dryas boundary sediment layer. *Science* 323:94–94.
- Kennett DJ, Kennett JP, West A, West GJ, Bunch TE (2009) Shock-synthesized hexagonal diamonds in Younger Dryas boundary sediments. *Proc Natl Acad Sci USA* 106:12623–12628.
- Firestone RB, West A, Kennett JP, Becker L, Bunch TE (2007) Evidence for an extraterrestrial impact 12,900 years ago that contributed to the megafaunal extinctions and the Younger Dryas cooling. *Proc Natl Acad Sci USA* 104:16016–16021.
- Paquay FS, Goderis S, Ravizza G, Vanhaeck F, Boyd M (2009) Absence of geochemical evidence for an impact event at the Bölling–Allerød/Younger Dryas transition. *Proc Natl Acad Sci USA* 106:21505–21510.
- Surovell T, Holliday VT, Gingerich JAM, Ketrona C, Haynes CV (2009) An independent evaluation of the Younger Dryas extraterrestrial impact hypothesis. *Proc Natl Acad Sci USA* 106:18155–18158.
- Daulton TL, Pinter N, Scott AC (2010) No evidence of nanodiamonds in Younger Dryas sediments to support an impact event. *Proc Natl Acad Sci USA* 107:16043–16047.
- Yang ZQ, et al. (2008) TEM and Raman characterisation of diamond micro- and nanostructures in carbon spherules from upper soils. *Diamond & Related Materials* 17:937–943.
- Egerton RF (1996) *Electron Energy-Loss Spectroscopy in the Electron Microscope* (Plenum, New York).
- Dadsetani M, Titantah JT, Lamoen D (2010) Ab initio calculation of the energy-loss near-edge structure of some carbon allotropes: Comparison with n-diamond. *Diam Relat Mater* 19:73–77.
- Matsumoto S, Matsui Y (1983) Electron microscopic observation of diamond particles grown from the vapour phase. *J Mater Science* 18:1785–1793.
- Kimoto K, Nishida I (1967) Multiply twinned particles of F.C.C. metals produced by condensation in argon at low pressures. *J Phys Soc Jpn* 22:940–940.
- Brown TE, LeMay HE, Bursten BE (1997) *Chemistry the Central Science*, (Simon & Schuster, Upper Saddle River, NJ), 227, pp 412–413.
- Banhart F, Pulickel M, Ajayan PM (1997) Self-compression and diamond formation in carbon onions. *Adv Mater* 9:261–263.
- Banhart F, Ajayan PM (1996) Carbon onions as nanoscopic pressure cells for diamond formation. *Nature* 382:433–435.
- Daulton TL, Eisenhour DD, Bernatowicz TJ, Lewis RS, Buseck PR (1996) Genesis of pre-solar diamonds: Comparative high-resolution transmission electron microscopy study of meteoritic and terrestrial nano-diamonds. *Geochim Cosmochim Acta* 60:4853–4872.
- Goresy EL (1968) A new allotropic form of carbon from the Ries crater. *Science* 161:363–364.
- Blake DF, Freund F, Krishnan KFM, Echer CJ, Shipp R (1988) The nature and origin of interstellar diamond. *Nature* 332:611–613.
- Lewis RS, Tang M, Wacker JF, Anders E, Steel E (1987) Interstellar diamonds in meteorites. *Nature* 326:160–162.
- Erskine DJ, Nellis WJ (1991) Shock-induced martensitic phase transformation of oriented graphite to diamond. *Nature* 349:317–319.
- Turner S, et al. (2009) Determination of size, morphology, and nitrogen impurity location in treated detonation nanodiamond by transmission electron microscopy. *Adv Funct Mater* 19:2116–2124.
- Hata T, Imamura Y, Kobayashi E, Yamane K, Kikuchi K (2000) Onion-like graphitic particles observed in wood charcoal. *J Wood Sci* 46:89–92.
- Cohen-Ofri I, Popovitz-Biro R, Weiner S (2007) Structural characterization of modern and fossilized charcoal produced in natural fires as determined by using electron energy loss spectroscopy. *Chem Eur J* 13:2306–2310.
- Blank VD, Kulnitskiy BA, Perezhogin IA (2009) Structural peculiarities of carbon onions, formed by four different methods: Onions and diamonds, alternative products of graphite high-pressure treatment. *Scripta Mater* 60:407–410.
- Blank VD, Kulnitskiy BA, Dubitskiy GA, Alexandrou I (2005) The structures of C60-Phases, formed by thermobaric treatment: HREM-studies. *Fuller Nanotub Car N* 13(supp. 1):167–177.
- Sano N, et al. (2002) Properties of carbon onions produced by an arc discharge in water. *J Appl Phys* 92:2783–2788.
- Sano N, Wang H, Chhowalla M, Alexandrou I, Amaratunga GAJ (2001) Nanotechnology: Synthesis of carbon “onions” in water. *Nature* 414:506–507.
- Chhowalla M, et al. (2003) Carbon onions: Carriers of the 217.5 nm interstellar absorption feature. *Phys Rev Lett* 90:155504–1.
- Cami J, Bernard-Salas J, Peeters E, Malek SE (2010) Detection of C60 and C70 in a young planetary nebula. *Science* 329:1180–1182.
- Hoefs J (2009) *Stable Isotope Geochemistry* (Springer, New York), 6th ed., p 288.
- Cartigny P (2005) Stable isotopes and the origin of diamond. *Elements* 1:79–84.
- Meyers PhA (1994) Preservation of elemental and isotopic source identification of sedimentary organic matter. *Chem Geol* 114:289–302.
- Russell SS, Arden JW, Pillinger CT (1996) A carbon and nitrogen isotope study of diamond from primitive chondrites. *Meteorit Planet Sci* 31:343–355.
- Pearson VK, Septhon MA, Franchi IA, Gibson JM, Gilmour I (2006) Carbon and nitrogen in carbonaceous chondrites: Elemental abundances and stable isotopic compositions. *Meteorit Planet Sci* 41:1899–1918.
- Alexander CMOD, Fogel M, Yabuta H, Gody GD (2007) The origin and evolution of chondrites recorded in the elemental and isotopic compositions of their macromolecular organic matter. *Geochim Cosmochim Acta* 71:4380–4403.
- McKeegan KD, Aléon J, Bradley J, Brownlee D, Busemann H (2006) Isotopic compositions of cometary matter returned by stardust. *Science* 314:1724–1728.
- Claeys Ph, Paquay F, Goderis S, Vanhaeck F (2008) Do the concentrations of platinum group elements in the Younger Dryas black layer really support an extraterrestrial origin? *American Geophysical Union, Fall Meeting 2008 Abstract* 89(53):P31A–1382.
- Koeberl C, Peucker-Ehrenbrink B, Reimold U, Shukolyukov A, Lugmair G (2002) Comparison of the osmium and chromium isotopic methods in the detection of meteoritic components in impactites: Examples from the Morokweng and Vredefort impact structure South Africa. *Geological Society of America, Special Publication* 356:607–617.

Materials and Methods

The material was collected on-site by one of the authors (D.S.) and small parts of the excavated material was crushed in ethanol and dispersed on holey carbon TEM grids. High-resolution TEM imaging as well as electron energy loss spectroscopy were performed with an FEI CM30 ultratwin field emission gun instrument operating at 300 keV equipped with a postcolumn GIF200 spectrometer. Background subtraction and zero-loss convolution of the spectra were obtained via the GATAN DM software.

Carbon and nitrogen content along with carbon isotope signature ($\delta^{13}\text{C}$) were determined by combusting preweighed crushed samples, acidified to remove CaCO_3 , and contained in silver cups in a Flash1112 elemental analyzer coupled to a Delta + XL via a conflat III interface (Thermo Scientific). Internal reference materials included IAEA-CH6 for C whereas acetanilide was used for total organic carbon and total nitrogen calibration. Typical reproducibility of $\delta^{13}\text{C}$ is $\pm 0.2\text{‰}$.